

# Phytoplankton photosynthetic parameters from the Gulf of California southern region

Parámetros fotosintéticos del fitoplancton en la región sur del Golfo de California

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**Resumen.-** Se determinó la respuesta fisiológica del fitoplancton en la zona eufótica por medio de experimentos fotosíntesis-irradiancia realizada de junio del 2000 a junio del 2001 en la Bahía de La Paz, Golfo de California, México. Durante el periodo estratificado que comprende de junio a septiembre (índice de estratificación  $> 200 \text{ J m}^{-3}$ ) el número de asimilación ( $P_m$ ) estuvo representado por un valor promedio de  $3,7 \text{ mg C m}^{-3} \text{ h}^{-1}$ , mientras que de noviembre a febrero, cuando la columna de agua fue homogénea (índice de estratificación =  $0 \text{ J m}^{-3}$ ) fue de  $9,2 \text{ mg C m}^{-3} \text{ h}^{-1}$ . La eficiencia fotosintética máxima a baja irradiancia ( $\alpha$ ) tuvo un valor promedio de  $0,017 \text{ mg C m}^{-3} \text{ h}^{-1} (\mu\text{mol quanta m}^{-2} \text{ s}^{-1})^{-1}$  representativo del periodo estratificado, mientras que durante el periodo homogéneo fue de  $0,063 \text{ mg C m}^{-3} \text{ h}^{-1} (\mu\text{mol quanta m}^{-2} \text{ s}^{-1})^{-1}$ . Los parámetros fotosintéticos ( $P_m$  y  $\alpha$ ) durante la mayor dinámica de la zona epipelágica evidencian la fotoaclimatación del fitoplancton a la irradiancia promedio en la zona eufótica. Valores altos de  $P_m$  se relacionaron con la abundancia de diatomeas (*Pseudo-nitzschia delicatissima*, *Chaetoceros compressus* y *Coscinodiscus perforatus*), mientras que los valores menores de este parámetro se asociaron con el grupo de los dinoflagelados. Los valores promedio estimados para el parámetro de saturación de luz ( $E_k$ ) fueron 334, 173 y 91  $\mu\text{mol quanta m}^{-2} \text{ s}^{-1}$  para superficie, 10 y 1% de la zona eufótica, respectivamente. De manera general, la comunidad fitoplanctónica estuvo dominada por nanofitoplancton (2-20  $\mu\text{m}$ ) durante todo el periodo considerado.

**Palabras clave:** Parámetros fotosintéticos, zona eufótica, Golfo de California

**Abstract.-** Phytoplankton physiologic response in the euphotic zone was determined by photosynthesis-irradiance experiments carried out between June 2000 and June 2001 at Bahia de La Paz, Gulf of California, Mexico. A stratified water column was present, from June to September (stratification index  $> 200 \text{ J m}^{-3}$ ), with the assimilation number ( $P_m$ ) represented by an average of  $3.7 \text{ mg C m}^{-3} \text{ h}^{-1}$ , while for the period from November through February, the water column was homogeneous (stratification index =  $0 \text{ J m}^{-3}$ ) and the average was  $9.2 \text{ mg C m}^{-3} \text{ h}^{-1}$ . The maximum photosynthetic efficiency at low irradiances ( $\alpha$ ) of  $0.017 \text{ mg C m}^{-3} \text{ h}^{-1} (\mu\text{mol quanta m}^{-2} \text{ s}^{-1})^{-1}$  was the representative average for the stratified period, while  $0.063 \text{ mg C m}^{-3} \text{ h}^{-1} (\mu\text{mol quanta m}^{-2} \text{ s}^{-1})^{-1}$  was the corresponding value for the homogeneous period. The photosynthetic parameters ( $P_m$  and  $\alpha$ ) had an increase during mixing conditions, as result of phytoplankton photoacclimation to average irradiance in the euphotic zone. Diatom abundance (*Pseudo-nitzschia delicatissima*, *Chaetoceros compressus* and *Coscinodiscus perforatus*) were related with higher values of  $P_m$ , while lowest values of this parameter are associated with the dinoflagellate group. Average values estimated for the light saturation parameter ( $E_k$ ) were 334, 173 and 91  $\mu\text{mol quanta m}^{-2} \text{ s}^{-1}$  for the surface, 10 and 1% of the euphotic zone, respectively. In general, nanophytoplankton (2-20  $\mu\text{m}$ ) dominated the phytoplanktonic community during the period studied.

**Key words:** Photosynthetic parameters, euphotic zone, Gulf of California

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## INTRODUCTION

One of the biological oceanography goals is the determination of the photosynthetic or carbon assimilation rate by phytoplankton (Falkowsky & Raven

1997). A perspicuous understanding of this complex process that controls marine biogeochemistry is essential, since their interlinked effects span from a regional to a

global scale. One investigative technique utilized to study this process is the determination of phytoplankton photosynthetic parameters from photosynthesis irradiance (P-E) curves (Jassby & Platt 1976). These allow the physiological response of the phytoplankton community in the face of environmental variability to be quantified (Coté & Platt 1983, 1984). The photosynthetic parameters  $\alpha$  (maximum light utilization coefficient at low irradiance) and  $P_m$  (maximum photosynthetic rate at light saturation), are related by the following equation:  $E_k = P_m / \alpha$ , where  $E_k$  is the light-saturation parameter ( $\mu\text{mol quanta m}^{-2} \text{ s}^{-1}$ );  $P_m$  is the maximum photosynthesis at light saturation ( $\text{mg C m}^{-3} \text{ h}^{-1}$ ) and  $\alpha$  is the maximum light utilization coefficient [ $\text{mg C m}^{-3} \text{ h}^{-1} (\mu\text{mol quanta m}^{-2} \text{ s}^{-1})^{-1}$ ]. These 2 parameters are fundamental to describe photosynthesis-irradiance relationships (Tilzer *et al.* 1993, Sakshaug *et al.* 1997).

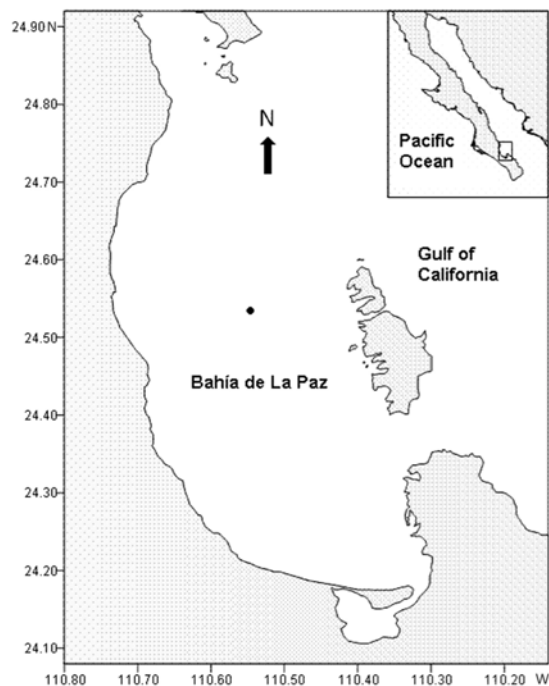
Phytoplankton photosynthetic parameters variability can be attributed to changes in environmental or growth conditions (Dower & Lucas 1993), and its delineation can ultimately lead to a better understanding of algal physiology and water column production. In this context, it is necessary to generate 'local specific' data for these photosynthetic parameters. Bahía de La Paz is a unique region due to its climatic regime for the set up and development of algorithms to estimate integrated water column production using phytoplankton P-E parameters. In addition, minimal yearly cloud cover favors the use of satellite color imagery which when combined with these parameters allows primary production to be estimated from satellite data (Platt & Sathyendranath 1988).

Estimates of primary productivity are necessary to assess the interchange of material comprising the ocean-atmosphere  $\text{CO}_2$  balance. Primary productivity data is sparse for this area of the Gulf of California, and when available it only documents indirect measurements of the photosynthetic rate (*i.e.*, phytoplankton natural fluorescence) (Reyes-Salinas *et al.* 2003, Cervantes-Duarte *et al.* 2005). Thus, due to the previously noted conditions this study will utilize the variability of photosynthesis-irradiance parameters of phytoplankton in Bahía de La Paz, with the aim of proposing representative average values for different oceanographic conditions, which in turn will allow prediction of photosynthetic production rates (Alvarez-Borrego 2011) based on models like those proposed by Platt & Sathyendranath (1988).

## MATERIALS AND METHODS

Bahía de La Paz is located in the southeast of the Baja California peninsula, bounded by lines making a polygon connecting the following geographical positions,  $24^\circ 10' - 24^\circ 47' \text{N}$ , and  $110^\circ 20' - 110^\circ 44' \text{W}$ , covering an approximate area of  $2,635 \text{ km}^2$  (Cruz-Orozco *et al.* 1989). The bay has a semi-elliptical shape, with its major axis oriented in a northwesterly to southeasterly direction. The bay communicates with the Gulf of California by 2 channels, a northeastern channel (Boca Grande) and another one to the east (Canal de San Lorenzo). During the warm period (May to October), intense solar radiation and the weak sea breezes from the southeast allow a strong stratification in the water column. During the cold season (November to April), strong northwest winds increase the mixed layer depth.

Monthly samplings were conducted at a station ( $24^\circ 29.12' \text{N}$ ,  $110^\circ 28.33' \text{W}$ ) located in the central area of Bahía de La Paz, (Fig.1) from June 2000 to June 2001, except July and December 2000 (Table 1). Samples were collected from 6 depths differentiated by the percentages



**Figure 1. Study area and localization of the sampling station (●) in the southern Gulf of California, México / Área de estudio y localización de la estación de muestreo (●) en el sur del Golfo de California, México**

**Table 1. Sampling dates and depths (m) of the surface irradiances percentages / Fechas de muestreo y profundidad (m) de los porcentajes de luz referidos a la irradiancia superficial**

Date (DD/MM/YY)	Surface irradiance percentages		
	100%	10%	1%
26/06/00	0	10	20
10/08/00	0	30	60
7/09/00	0	29	58
16/10/00	0	27	55
26/11/00	0	20	40
26/01/01	0	20	40
14/02/01	0	20	40
22/03/01	0	17	36
20/04/01	0	20	40
21/05/01	0	14	29
22/06/01	0	14	29

of light penetration (100, 55, 33, 10, 3 and 1% of surface irradiance), calculated using the light attenuation coefficient ( $K_d$ ) determined by the Secchi disk (Brown *et al.* 1989). Sampling depths were estimated following Lambert-Beer relationship  $Z = (Ln E_o - Ln E_z) / K_d$ . At each depth, water column was collected for taxonomic phytoplankton identification, and dissolved inorganic nutrients analyses. Phytoplankton samples were fixed with acid Lugol's solution (final concentration of 1%), and analyzed employing the method devised by Utermöhl (Hasle 1978). The phytoplankton taxa were identified according to Hustedt (1930, 1959), Cupp (1943), Licéa-Durán (1974) for diatoms; Schiller (1933), Taylor (1976), Pesantes (1978) and Balech (1988) for dinoflagellates and Murray & Schrader (1983) and Schrader *et al.* (1986), for silicoflagellates. Hasle & Syvertsen (1996) and Steindinger & Tangen (1996) were consulted to review and update in cases of synonymy between species.

The concentration of inorganic nutrients (nitrites and nitrates) was determined according to the methods proposed by Parsons *et al.* (1984). The properties of the water column were characterized (temperature, salinity) using a CTD (Sea Bird-25). The stratification index proposed by Simpson & Hunter (1974) was calculated using the temperature, salinity and density data for the upper 50 m depth. The stratification index ( $\phi$ ), interpreted as the amount of mechanical energy required to bring about vertical mixing ( $J m^{-3}$ ), equals zero for a well-mixed layer and increases with stratification, expressed by the following equation:

$$\phi = \frac{1}{h} \int_{-h}^0 (\bar{\rho} - \rho) g z dz$$

where:  $\rho$  = seawater density,  $\bar{\rho}$  = average seawater density,  $g$  = gravity acceleration ( $9.81 m s^{-2}$ ),  $h$  = depth (in this case 50 m), and  $z$  = vertical coordinate (ascendent and positive).

The determination of photosynthetic parameters was carried out by experimental measurements of  $^{14}C$  assimilation during June, September and November of 2000, as well as during February and May of 2001. Water samples were collected from 100%, 10% and 1% surface irradiance levels, using dark bottles and filtered through a 333  $\mu m$  mesh to remove microzooplankton. Samples were stored in refrigerator while the experiment was prepared. Water samples were collected with a lapse of approximately 1-1.5 h between sample acquisition and the initiation of incubation in the laboratory, where samples were inoculated with 1  $\mu Ci$  of  $NaH^{14}CO_3$ . Inoculated samples were incubated during 1 h in 3 incubators similar to those proposed by Babin *et al.* (1994). Incubation bottles were the same type as those used for tissue cultures, polystyrene plastic, which have provided good results in the incubation of marine phytoplankton samples; polystyrene doesn't affect the phytoplankton for incubation periods up to 2 h (Babin *et al.* 1994). The light source was a 300 W tungsten-halogen lamp. Care was taken during the incubation process, so that water temperature inside incubators remained approximately equal to the temperature from which the sample was collected ( $\pm 1^\circ C$ ). The radioactivity counts in the phytoplankton samples were measured utilizing a Beckman LSC 6500 liquid scintillation counter. The inorganic carbon assimilation was determined following Steemann-Nielsen (1952).

Mathematical model used for the adjustment of the photosynthetic parameters was the proposed by Jassby & Platt (1976). Nonlinear models from Statistica ver. 7 software were used to simultaneously estimate both parameters ( $P_m$  and  $\alpha$ ) using the relationship:

$$P = P_m \tanh(\alpha E_o / P_m)$$

where  $\alpha$  is the maximum light coefficient at low irradiance [ $(mg C m^{-3} h^{-1}) (\mu mol quanta m^{-2} h^{-1})^{-1}$ ],  $P_m$  is the maximum photosynthesis at light saturation ( $mg C m^{-3} h^{-1}$ ),  $E_o$  is the light intensity in the incubator ( $\mu mol quanta m^{-2} s^{-1}$ ).

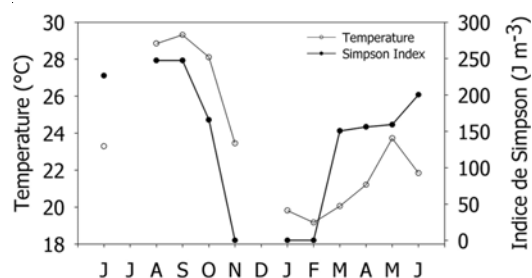
## RESULTS

Water column temperature and the Simpson & Hunter (1974) stratification index ( $\phi$ ;  $J m^{-3}$ ) indicated stratification from June through September. This period was characterized by temperature in excess of  $28^{\circ}C$ , and high stratification index ( $>200 J m^{-3}$ ); while from November to February exhibited a mixed layer depth exceeding 50 m deep (*i.e.*, zero stratification) and lower temperature. These conditions began to reversal slowly, rising trend starting in March, until the stratification index reached  $200 J m^{-3}$  in June 2001. From this parameter May 2001 was considered a transitional month (Fig. 2).

Phytoplankton abundance was dominated both temporally and spatially by nanophytoplankton (cell size 2-20  $\mu m$ ), with higher cells abundance recorded at subsurface levels during June and September 2000. The microphytoplankton fraction (cells size  $>20 \mu m$ ) showed their maximum proliferation in June 2000 at the bottom of the euphotic zone (Table 2), with *Pseudo-nitzschia delicatissima* as the the dominant specie.

Nitrate concentration in the euphotic zone showed a stratified distribution all year. Minimal values ( $<2.5 \mu M$ ) were associated with the surface layer, mainly during the months of higher water column stratification (from July

to September 2000); while maximum values were registered at the deepest levels of the euphotic zone. We observed a nitrate increase at 10 and 1% irradiance levels mainly during October 2000 and March 2001. Nitrite concentrations in the euphotic zone registered values lower than  $0.2 \mu M$  from June to October 2000. Early in the condition of mixed water column, nitrite concentration gradually increased until its maximum value achieved in March ( $1.2 \mu M$ ). Subsequently, a decrease was observed during April, May, and June (Table 2).



**Figure 2.** Monthly sea surface temperature and the Simpson and Hunter (1974) stratification index from June 2000 to June 2001 / Variabilidad mensual de la temperatura superficial del mar y el índice de estratificación de Simpson y Hunter (1974) de junio de 2000 a junio de 2001

**Table 2.** Nano- (5-20  $\mu m$ ) and microphytoplankton ( $>20 \mu m$ ) abundance, and nutrient concentration / Abundancia del nano (5-20  $\mu m$ ) y microfitoplancton ( $>20 \mu m$ ) y concentración de nutrientes

Year/Month	Nanophytoplankton ( $10^3$ cells $L^{-1}$ )			Microphytoplankton ( $10^3$ cells $L^{-1}$ )			Nitrite ( $\mu M$ )			Nitrate ( $\mu M$ )		
	100%	10%	1%	100%	10%	1%	100%	10%	1%	100%	10%	1%
2000												
June	242.2	575	201	1.4	6.2	14.40	0.010	0.010	0.020	0.480	19.08	13.10
August	228	245	210	1.6	0.50	1	0.006	0.024	0.012	0.421	3.16	10.22
September	301	546.2	410	1.1	1.7	0.80	ND	0.040	0.020	0.510	3.51	7.38
October	456	512	486	0.5	1.3	1	0.006	0.059	0.008	0.279	10.86	18.65
November	385	402	268	1.5	1.3	1.90	0.380	0.480	0.440	3.060	2.77	4.88
2001												
January	326.2	401	298.3	1.5	1	0.90	0.059	0.054	0.071	2.630	2.14	3.55
February	452.3	488	399.9	1.2	1.2	1	0.030	0.740	0.660	22.26	5.70	15.18
March	102.8	99.2	86.4	0.9	1.1	0.40	0.060	0.042	0.063	0.360	4.03	10.01
April	118	108.6	99	2.6	4	4.80	ND	ND	ND	ND	ND	ND
May	62	61.6	52	9.3	6.3	4.10	ND	ND	ND	ND	ND	ND
June	206	196	196.6	0.8	1.8	1	0.050	0.032	0.45	0.390	13.4	7.1

\*ND= no data available

Phytoplankton photosynthetic parameters showed a temporal variability related with oceanographic conditions of the water column. During months of stratification (June and September 2000), maximum photosynthetic efficiency at low irradiances ( $\alpha$ ) from the 3 levels of the sampled euphotic zone fluctuating between 0.005 and 0.035 mg C m<sup>-3</sup> h<sup>-1</sup> ( $\mu\text{mol quanta m}^{-2} \text{s}^{-1}$ )<sup>-1</sup>, giving a representative mean of 0.017 mg C m<sup>-3</sup> h<sup>-1</sup> ( $\mu\text{mol quanta m}^{-2} \text{s}^{-1}$ )<sup>-1</sup>. During November 2000 and 2001, when the water column was homogeneous, ( $\phi \sim 0 \text{ J m}^{-3}$ ),  $\alpha$  reached their maxima values [0.052 and 0.115 mg C m<sup>-3</sup> h<sup>-1</sup> ( $\mu\text{mol quanta m}^{-2} \text{s}^{-1}$ )<sup>-1</sup>], and the difference between the 3 sampled levels increase, showing a variation between 0.019 and 0.115 mg C m<sup>-3</sup> h<sup>-1</sup> ( $\mu\text{mol quanta m}^{-2} \text{s}^{-1}$ )<sup>-1</sup>. The overall representative average was 0.06 mg C m<sup>-3</sup> h<sup>-1</sup> ( $\mu\text{mol quanta m}^{-2} \text{s}^{-1}$ )<sup>-1</sup>. During May considered as a transitional month between homogeneity and stratification conditions,  $\alpha$  ranged from 0.005 and 0.007, with a mean value of 0.006 mg C m<sup>-3</sup> h<sup>-1</sup> ( $\mu\text{mol quanta m}^{-2} \text{s}^{-1}$ )<sup>-1</sup>.

Maximum photosynthesis rate ( $P_m$ ) behaved similarly to the previously described parameter. However, its variability between irradiance levels showed more deviation. For the stratification period,  $P_m$  ranged from 0.72 to 8.66 mg C m<sup>-3</sup> h<sup>-1</sup>, with a mean of 3.62 mg C m<sup>-3</sup> h<sup>-1</sup>. In contrast, this parameter showed markedly higher values during the months of mixed water column ( $\phi \sim 0 \text{ J m}^{-3}$ ), ranging between 6.39 and 12.69 mg C m<sup>-3</sup> h<sup>-1</sup>, with a representative mean value of 9.3 mg C m<sup>-3</sup> h<sup>-1</sup>. During May 2001 the minimum registered value was 0.41 and the maximum 1.72, with an average of 1.14 mg C m<sup>-3</sup> h<sup>-1</sup>.

With respect to the light saturation index ( $E_k$ ), the main source of variation was on the vertical coordinate, as this

parameter showed evident differences between the considered light levels. Light saturation index at surface registered values between 232 y 469  $\mu\text{mol quanta m}^{-2} \text{s}^{-1}$ , with an average of 335  $\mu\text{mol quanta m}^{-2} \text{s}^{-1}$ . In the intermediate level (10% of surface irradiance), values oscillated between 110 and 247  $\mu\text{mol quanta m}^{-2} \text{s}^{-1}$ , with an average of 173  $\mu\text{mol quanta m}^{-2} \text{s}^{-1}$ . In the lower limit of the euphotic zone (1% of surface irradiance)  $E_k$  changed between 65 and 119  $\mu\text{mol quanta m}^{-2} \text{s}^{-1}$ , being 91  $\mu\text{mol quanta m}^{-2} \text{s}^{-1}$  the average value (Table 3). Euphotic zone lowest differences of  $E_k$  were obtained during November and February (water column mixing conditions), indicating highest phytoplankton photoacclimation to mean irradiance, derived from the lower stratification prevailing during those months.

By correlating the Simpson stratification index ( $\phi$ ) with the photosynthetic parameters, we found that the maximum photosynthetic efficiency at low irradiances ( $\alpha$ ) showed negative slope for the 3 considered light levels, with the determination coefficient ( $r^2$ ) between 0.54 and 0.72. In the same sense, maximum photosynthetic parameter ( $P_m$ ) had a negative linear relationship, with  $\phi$  showing a variation between 0.45 and 0.81 in the determination coefficient. Unlike the 2 previously described parameters, light saturation index ( $E_k$ ) showed positive slope with  $\phi$  at both the surface and the level corresponding to 10% of surface irradiance, with determination coefficients of 0.46 and 0.81, respectively. At the lower limit of the euphotic zone, the determination coefficient was higher (0.91), with negative slope.

**Table 3. Phytoplankton photosynthetic parameters at 100, 10 and 1% of surface irradiance / Parámetros fotosintéticos del fitoplancton a 100, 10 y 1% de la irradiancia superficial**

Month	Photosynthetic parameters								
	$P_m$ mg C m <sup>-3</sup> h <sup>-1</sup>			$\alpha$ [mg C m <sup>-3</sup> h <sup>-1</sup> ( $\mu\text{mol quanta m}^{-2} \text{s}^{-1}$ ) <sup>-1</sup> ]			$E_k$ $\mu\text{mol quanta m}^{-2} \text{s}^{-1}$		
	100%	10%	1%	100%	10%	1%	100%	10%	1%
June	1.75	2.93	0.72	0.005	0.016	0.011	350	183	65
September	5.63	8.66	2.02	0.012	0.035	0.026	469	247	78
November	12.06	10.90	7.02	0.052	0.078	0.059	232	139	119
February	6.39	12.69	6.59	0.019	0.115	0.060	336	110	109
May	1.72	1.29	0.41	0.006	0.007	0.005	288	184	82



## DISCUSSION

Phytoplankton studies have been carried out during the last decade in the 2 larger bays of the Gulf of California (Bahía de La Paz and Bahía Concepción). These studies have documented that primary productivity and phytoplankton community structure display seasonal changes, associated mainly with the occurrence of strong stratification during summer, and upwelling and water column mixing during winter (Martínez-López *et al.* 2001, Verdugo-Díaz *et al.* 2010). Biological responses to the physical forcing suggest that changes in phytoplankton community structure, as evidenced by increases of abundance in nanophytoplankton during summer conditions, could be related to a greater capacity of these cells to acquire and assimilate lower nutrient concentration during oligotrophic-water periods; allowing them to obtain a competitive advantage over larger cells (Fogg *et al.* 1986). Another factor, which possibly contributes to their dominance throughout the water column, may be related to a greater efficiency in their absorption and utilization of incident light compared to larger-sized cells, particularly at low irradiances (Agustí *et al.* 1994).

Photosynthesis-irradiance related parameters contain information of physiological seasonal changes of the phytoplankton community in Bahía de La Paz. These parameters according with their characteristics can be applied to specific examples and utilized to generate estimates of primary production in the study zone (Valdez-Holguín *et al.* 1999). Phytoplankton photosynthesis rate variability in the ocean mainly depends of the quantity and quality of available light for the microalgae growth. Walsh & Legendre (1983) did not notice significant differences in phytoplankton assimilation rates produced by variable irradiance, because photosynthetic capacity is not a direct function of light absorption (Sakshaug *et al.* 1997), it also depends on photo-acclimation. Marine microalgae exhibit several irradiance requirements in order to saturate their photosynthetic process and achieve their maximum productivity.

The P-E curve derived parameters generated in this study are high when compared to those reported for the central region of the Gulf of California (Gaxiola-Castro *et al.* 1999, Valdéz-Holguín *et al.* 1999), but similar to values registered around the large islands (Tiburón & Ángel de la Guarda) in the northern Gulf (Alvarez-Borrego & Gaxiola-Castro 1988) as well as for a subtropical coastal lagoon of the Baja California Peninsula (Montes-Hugo & Alvarez-Borrego 2003). It seems that oceanographic conditions at

Bahía La Paz support high values of primary and secondary productivity (Jiménez-Illescas 1994, Martínez-López *et al.* 2001, De Silva-Dávila & Palomares-García 2002), making it one of the most productive zones of the Gulf.

Nutrient concentrations recorded in this study are consistent with those reported by Reyes-Salinas *et al.* (2003) and Cervantes-Duarte *et al.* (2005) for Bahía de La Paz. They described low nitrate concentrations in the upper euphotic zone during the entire year, which declined during the summer period. In our study we found high correlation between the values of  $P_m$  and nitrate concentration (data not shown). Gaxiola-Castro *et al.* (1999) established that an increase in nutrient concentration is not necessarily correlated with an increase in  $P_m$ , because changes of this photosynthetic parameter are also a function of variations in growth irradiance controlled by stratification or mixing in the water column.

The observed values of maximum photosynthesis showed low variability during the studied period. Despite the low frequency of obtained data, which limits us to contrast statistically, it is clear that conditions of the water column play a central role in the magnitude of photosynthesis rate. Because of this, we consider very useful a mean value of  $3.6 \text{ mg C m}^{-3} \text{ h}^{-1}$  as representative of stratified conditions,  $9.3 \text{ mg C m}^{-3} \text{ h}^{-1}$  for homogeneity conditions, and  $1.1 \text{ mg C m}^{-3} \text{ h}^{-1}$  for the transitional periods. The lowest observed variability, especially during the stratified water column, may be associated to acclimation or tolerance ranges of phytoplankton in the face of environmental fluctuations that modify its photosynthetic parameters (Coté & Platt 1983). Besides irradiance, chlorophyll cell-quantity or concentration (first order factor) in phytoplankton community is considered as the most influential in the variability of parameters derived from P-E curves (Cullen *et al.* 1992, Marañón & Holligan 1999).

Under the same scheme, for the results corresponding to maximum photosynthetic efficiency at low irradiances, we suggest a value of  $0.017 \text{ mg C m}^{-3} \text{ h}^{-1}$  ( $\mu\text{mol quanta m}^{-2} \text{ s}^{-1}$ )<sup>-1</sup> as representative of stratification conditions; and an average value when the water column is homogeneous of  $0.06 \text{ mg C m}^{-3} \text{ h}^{-1}$  ( $\mu\text{mol quanta m}^{-2} \text{ s}^{-1}$ )<sup>-1</sup>. As the change is gradual between both conditions, the value that represents the transition period is  $0.006 \text{ mg C m}^{-3} \text{ h}^{-1}$  ( $\mu\text{mol quanta m}^{-2} \text{ s}^{-1}$ )<sup>-1</sup>. The understanding of variability of this phytoplankton physiological indicator

is important because it represents the trend of phytoplankton towards equilibrium between the light phase and dark reactions of photosynthesis; like equilibrium between quanta absorption for photosystem II and the electrons transported from water in order to reduce carbon dioxide (Escoubas *et al.* 1995, Sakshaug *et al.* 1997). Regarding this, Tilzer (1993) discussed the importance of understanding its magnitude and variability, as he defined it as a function of the number of photosynthetic units or reaction centers and antenna pigments.

Light saturation index parameter ( $E_k$ ) showed a marked correlation with depth, while the extent of temporal variation was lower. Considering these observations, we suggest an average of  $335 \mu\text{mol quanta m}^{-2} \text{s}^{-1}$  as the irradiance that saturates the phytoplankton photosynthetic process at ocean surface. Despite the relatively low value of  $E_k$  in this case, we did not observe photoinhibition—a decline process in quantum efficiency caused by both irradiance and increasing temperature (Anderson *et al.* 1997). It has been observed that in natural environments like the Gulf of California, surface phytoplankton shows no photoinhibition, even at irradiances as high as  $1.400 \mu\text{mol quanta m}^{-2} \text{s}^{-1}$  (Valdez-Holguín *et al.* 1999).

From phytoplankton samples collected at 10% and 1% surface irradiance levels, maxima photosynthesis reach saturation at  $173.2 \mu\text{mol quanta m}^{-2} \text{s}^{-1}$  and  $90.8 \mu\text{mol quanta m}^{-2} \text{s}^{-1}$ , respectively. During each stratified period, a difference between levels was observed. This could be related with density differences which arguably propitiate that phytoplankton cells remain at the same depth, which in turn may allow better cell photoacclimation (Ley 1980, Falkowski 1983). On the other hand, during months of homogeneous water column, phytoplankton communities at the edge of euphotic zone may move upwards and acclimate accordingly to higher irradiances (Falkowski 1983), producing the observed uniformity between the  $E_k$  values corresponding to irradiance levels of 10 and 1% of surface irradiance. It is well known that phytoplankton photoacclimation processes occur within time intervals varying from hours to days (Ley 1980), according to requirements and pigment structure of the species existent in the communities (Bidigare *et al.* 1990, Falkowski & La Roche 1991). There are 2 main strategies for phytoplankton photoacclimation to dynamic irradiance regimes: 1) increments in photosynthetic unit's size, and 2) increase in the number of photosynthetic units. The physiologic mechanisms that occur at cellular level in adaptation to

low irradiances comprise both, with a decrease in the numbers of photosynthetic units, and increase in the cell chlorophyll-*a* (Ley 1980). In the same way, cells exposed to high photonic densities may resort to the xanthophylls cycle. Under such conditions, photon excess in the antenna unit is reflected as violaxanthin transformation into zeaxanthin helping to the dissipation of excessive energy (Hall 1994).

Hydrographic conditions at La Paz Bay maintain a stratified system from June to September, while in the period between November to February water-column remains homogeneous. Gaxiola-Castro *et al.* (1999) suggested that both, high water column stratification and mixing lead to low values in the water-column averaged photosynthetic parameters. Phytoplankton community react to this variability, with dinoflagellates and nanophytoplankton being more abundant during stratified conditions, while diatoms reach greater occurrence during the months of homogeneous water column (data not shown). In the same approach, the photosynthetic parameters as indicators of phytoplankton physiologic response in the expression of environmental variability show a relationship between its values and the prevailing environmental conditions. Despite the low frequency of experimental data, this study provides important representative values for  $P_m$  and  $\alpha$  at different oceanographic conditions in the water column. Likewise, representative values of  $E_k$  were proposed for phytoplankton sampled communities at different irradiance levels. Our approach is based on backgrounds suggesting a closed relationship between water column physical stability and photosynthetic parameters for the southern Gulf of California phytoplankton.

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